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SPH-Based Simulations for Slope Failure Considering Soil-Rock Interaction

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Abstract

Both small and large deformations of soil and soil-rock interaction are the key features in the post-failure process of natural slopes. As one of Lagrange meshless particle methods, the Smooth Particle Hydrodynamics (SPH) method has obvious advantage in dealing with the large deformation and interface interaction problems. Thus in this paper, the SPH method is employed to study the slope failure problems, especially focusing on soil large deformation and soil-rock interaction in natural slopes constituted of earth-rock aggregate. The Drucker-Prager model is implemented into the SPH-code to describe the elastic-plastic soil behavior while the rocks are simulated as rigid bodies by using classical rigid motion equations. The interaction between soil and rocks is modeled by the coupling condition associated with an action and reaction force between the two phases. Rock-rock contacts are computed using contact mechanics theory which is similar to the treatment in the Discrete Element Method (DEM). Two test cases including uniform non-cohesive and cohesive soil slopes failure problems are studied respectively to validate the method and a good agreement with the experimental data or numerical results is observed. Then the solution to soil-rock interaction is applied to study the behavior of earth-rock aggregate in the failure process of typical natural landslides. Numerical results show that the proposed soil-rock interaction algorithm works well in the SPH framework and has a great application potential in geotechnical engineering. Through the qualitative analysis and discussions about the behavior of earth-rock aggregate, we came to the conclusion that the soil-rock interaction has a great influence on the landslide shape in terms of rock size and slope angle. And the deformation characteristics of landslide at the slope toe is slightly different with that at the slope top.

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1. Introduction

The slope failure usually has the destructive potential both in terms of human lives and their property, as demonstrated by the landslide that hit Afghanistan May 2014. Therefore, the research on time prediction of slope failure is significant. As we known, natural slopes generally consist of earth-rock aggregate. In the event of landslide, the earth-rock aggregate easily produces fluidity and the soil inevitably interacts with the rocks. Thus it can be seen both small and large deformations of soil and soil-rock interaction are the key features in the post-failure process of natural slopes.

The earth-rock aggregate is a kind of special geological material between soil mass and rock mass, which has the non-linear property with obvious discontinuity, irregularity, and uncertainty. So it is difficult to utilize the traditional numerical method to make practical description and analyze its peculiar geological and mechanical behavior. For example, the finite element Method (FEM) may suffer from the grid distortion problem, which needs laborious work to overcome. The discrete element method (DEM) could be better in dealing with these large deformation and significant flow problems. But for the continuous granular material, DEM would not outperform FEM in calculation accuracy. In addition, the specification of DEM parameters is somewhat ambiguous, and reliable guidelines haven't yet to be clearly established^[1].

The Smooth Particle Hydrodynamics (SPH) method has several advantages in dealing with large deformation of continuum or dispersed material and the interface interaction problems. The SPH method is a purely Lagrange meshless method in which particles carry field variables such as mass, density, stress tensor, etc. and move with the material velocity^[2]. It can handle large deformation and post-failure very well due to its Lagrangian and adaptive nature; complex free surfaces are captured naturally without any special treatments; complex geometries and interface interaction can be treated relatively easily. Fortunately, Bui et al.^[1] has implemented elastic-plastic soil constitutive models in SPH code which available solved plastic soil behavior. Afterwards, Opez, Y.R.L. et al.^[3] proposed a dynamic refinement procedure to reduce the computational requirements of an elastic-plastic model to simulate non-cohesive soil. They demonstrated that SPH has good performance to simulate large deformation and post-failure problems such as slope failures, landslides, particle flows that are common in geotechnical engineering.

The motivation of this paper is to discuss that how the earth-rock aggregate affects the deformation characteristics of slope failure. The post-failure with soil-rock and rock-rock interaction involves both small and large deformations. Considering the advantages of SPH method, this paper exploited it to simulate slope failure process and the soil-rock interaction. To begin with, based on elastic-plastic soil model, we establish a soil-rock interaction procedure in the DualSphysics code^[4-6] (www.dual.sphysics.org), which allows to study more details about the behavior of the earth-rock aggregate. Then, two test cases including uniform non-cohesive and cohesive soil slope failures subjected to gravitational loading are simulated respectively to validate the method and good agreements with the experimental data or numerical results are observed. Finally, this method is employed to study the mechanics in failure process of typical natural slope constituted of earth-rock aggregate.

Nomenclature

| | |
|-----------------------------|---|
| α, β | Greek superscripts denoting Einstein's notation |
| i, j | Latin subscripts denoting individual particles |
| $f_g, \delta^{\alpha\beta}$ | gravitational acceleration and the Dirac delta function, respectively |
| m, ρ, v, x | mass, density, velocity and position of a soil particle, respectively |
| M, V, Ω | mass, velocity and rotational velocity of a rock, respectively |
| I, R_0 | the moment of inertia and the position of the centre of rock mass, respectively |
| σ, ε | the total stress and strain tensor, respectively |
| c, φ | the Coulomb's material constants: cohesion and internal friction angle |
| f, g | the yield and plastic potential function, respectively |
| E, ν | Young's modulus and Poisson's ratio, respectively |
| G, K | the shear modulus and the elastic bulk modulus, respectively |
| I_1, I_2 | the first and second invariants of the stress tensor, respectively |
| $W, h, \Delta d$ | the cubic spline kernel function, the smoothing length and the initial particle spacing, respectively |

| | |
|------------|---|
| f_{rs} | the force per unit mass exerted by soil particle s on rock particle r |
| SPs, RPs | soil particles and rock particles within smoothed kernel, respectively |

2. SPH formulation

In this section, we first formulate the basic governing equations for soil. Then, we briefly describe an elastic-plastic soil model based on the Drucker-Prager yield condition in the SPH framework. More details can be found in the original work by Bui et al.^[1]. Finally, we present the SPH implementation of soil-rock interaction.

2.1. Continuity and momentum equations

The continuity equation in SPH can be expressed as

$$\frac{D\rho_i}{Dt} = \sum_{j=1}^N m_j (v_i^\alpha - v_j^\alpha) \frac{\partial W_{ij}}{\partial x_i^\alpha} \quad (1)$$

The momentum equation can be described by the following SPH discretization:

$$\frac{Dv_i^\alpha}{Dt} = \sum_{j=1}^N m_j \left(\frac{\sigma_i^{\alpha\beta} + \sigma_j^{\alpha\beta}}{\rho_i \rho_j} - \Pi_{ij} \delta^{\alpha\beta} + F_{ij}^n R_{ij}^{\alpha\beta} \right) \frac{\partial W_{ij}}{\partial x_i^\beta} + f_g^\alpha \quad (2)$$

The artificial viscosity term Π is used to stabilize the numerical system. It is given by

$$\Pi_{ij} = \begin{cases} \frac{-\alpha c \mu_{ij} + \beta \mu_{ij}^2}{\bar{\rho}_{ij}} & v_{ij} \cdot x_{ij} < 0 \\ 0 & v_{ij} \cdot x_{ij} \geq 0 \end{cases}, \quad (3)$$

where $\mu_{ij} = \frac{h v_{ij} \cdot x_{ij}}{x_{ij}^2 + 0.01 h^2}$, $\bar{\rho}_{ij} = \frac{\rho_i + \rho_j}{2}$, α and β are constants respectively taken to be 0.1 and 0 here.

The artificial stress method refers to the term $F_{ij}^n R_{ij}^{\alpha\beta}$ is applied to reduce tensile instabilities. The treatment is completely the same as suggested in the article [1].

2.2. The Drucker-Prager soil model

For elastic-plastic materials, the strain rate tensor is normally composed of two parts, the elastic $\dot{\epsilon}_e^{\alpha\beta}$ and the plastic $\dot{\epsilon}_p^{\alpha\beta}$ strain rate tensors, resulting in

$$\dot{\epsilon}^{\alpha\beta} = \dot{\epsilon}_e^{\alpha\beta} + \dot{\epsilon}_p^{\alpha\beta} \quad (4)$$

The elastic term can be calculated by the generalized Hooke's law:

$$\dot{\epsilon}_e^{\alpha\beta} = \frac{\dot{s}^{\alpha\beta}}{2G} + \frac{1-2\nu}{3E} \dot{\sigma}^{\gamma\gamma} \delta^{\alpha\beta} \quad (5)$$

The plastic part can be computed by using the plastic flow rule:

$$\dot{\epsilon}_p^{\alpha\beta} = \dot{\lambda} \frac{\partial g}{\partial \sigma^{\alpha\beta}} \quad (6)$$

The Drucker-Prager yield condition is applied here to determine the soil plastic flow regime. Accordingly, the plastic deformation will occur only if the following yield criterion is satisfied

$$f(I_1, J_2) = \sqrt{J_2} + \alpha_\phi I_1 - k_c = 0 \quad (7)$$

α_ϕ and k_c are the Drucker-Prager's constants, in plane strain condition, are computed by

$$\alpha_\phi = \frac{\tan \phi}{\sqrt{9 + 12 \tan^2 \phi}} \text{ and } k_c = \frac{3c}{\sqrt{9 + 12 \tan^2 \phi}} \quad (8)$$

The non-associated plastic flow rule specifies the plastic potential function by

$$g(I_1, J_2) = \sqrt{J_2} + \alpha_\psi I_1 - \text{constant} \quad (9)$$

α_ψ is similar to α_ϕ , given by

$$\alpha_\psi = \frac{\tan \psi}{\sqrt{9 + 12 \tan^2 \psi}} \quad (10)$$

The rate of change of the plastic multiplier λ is obtained by solving

$$\dot{\lambda}_i = \frac{3\alpha_\phi K \dot{\epsilon}_i^{\gamma\gamma} + (G/\sqrt{J_2}) \mathbf{s}_i^{\alpha\beta} \dot{\epsilon}_i^{\alpha\beta}}{9\alpha_\phi \alpha_\psi K + G} \quad (11)$$

Finally, the stress-strain relationship, in particle approximation form, is given by

$$\frac{D\sigma_i^{\alpha\beta}}{Dt} = \sigma_i^{\alpha\gamma} \dot{\omega}^{\beta\gamma} + \sigma_i^{\gamma\beta} \dot{\omega}_i^{\alpha\gamma} + 2G \dot{\epsilon}_i^{\alpha\beta} + K \epsilon_i^{\gamma\gamma} \delta_i^{\alpha\beta} - \dot{\lambda}_i \left[3\alpha_\psi K \delta^{\alpha\beta} + \frac{G}{\sqrt{J_2}} \mathbf{s}_i^{\alpha\beta} \right] \quad (12)$$

Where, the two first terms are the results from the Jaumann's stress rate tensor, the third and fourth terms refer to the elastic behavior, and the last term in the equation relates to the plastic deformation.

2.3. Soil-rock interaction

In this paper, we treat the rocks as rigid bodies and use the Newton's equation for rigid body dynamics to describe the rock motion. The force on each rock particle is computed by summing up the contribution from all the surrounding soil particles within the smooth kernel, as shown in Fig.1. Hence, rock particle r experiences a force per unit mass given by

$$\mathbf{f}_r = \sum_{s \in SPs} \mathbf{f}_{rs} \quad (13)$$

By the principle of equal and opposite action and reaction, the force exerted by a soil particle on each rock particle is obtained by $m_r \mathbf{f}_r = -m_s \mathbf{f}_s$. Thus we can estimate the force exerted on the whole moving body actually only by computing the force, \mathbf{f}_s , exerted by a rock particle on a soil particle.

The rigid body dynamics including the translation and rotation equations are given by

$$M \frac{d\mathbf{V}}{dt} = \sum_{r \in RPs} m_r \mathbf{f}_r \text{ and } I \frac{d\boldsymbol{\Omega}}{dt} = \sum_{r \in RPs} m_r (\mathbf{r}_r - \mathbf{R}_0) \times \mathbf{f}_r \quad (14)$$

Each rock particle within the rigid body is moved in time by the velocity

$$\mathbf{u}_r = \mathbf{V} + \boldsymbol{\Omega} \times (\mathbf{r}_r - \mathbf{R}_0) \quad (15)$$

Besides, rock-rock contacts are computed by contact mechanics theory which is similar to the treatment in DEM^[7].

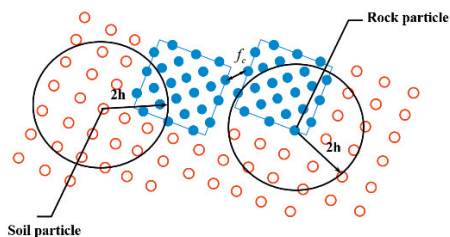


Fig.1. Sketch of soil-rock interaction.

Table 1. Soil parameters for test cases

| No. | Soil parameters | T-Case 1 | T-Case 2 |
|-----|-------------------|------------------------|------------------------|
| 1 | Density | 2650 kg/m ³ | 2100 kg/m ³ |
| 2 | Elastic modulus | 8.4 MPa | 15 MPa |
| 3 | Poisson ratio | 0.30 | 0.25 |
| 4 | Cohesive strength | 0 kPa | 11 kPa |

3.Validation and Verification

| | | | |
|---|-----------------|-------|-----|
| 5 | Friction angle | 19.8° | 20° |
| 6 | Dilatancy angle | 0° | 0° |

Two test cases are implemented here to validate the method. Fig.2. shows the initial state of the two test cases, respectively. The soil parameters are listed clearly in table 1. The first test (T-Case 1) is about uniform non-cohesive soil slope failure. We simulated the gravitational flow following soil collapse. Comparing the final shape between experiment and SPH simulation, a good agreement is observed as shown in Fig.3. The other one (T-Case 2) is about uniform cohesive soil slope failure. Comparing the process of slope failure at three representative time instants between Bui's and our simulated result, the results of the comparison are satisfied as shown in Fig.4. Both of the two test cases can indicate that large deformation of soil during the post-failure process can be described very well through the SPH simulation. And the SPH method has a great application potential in studying the behavior of earth-rock aggregate.

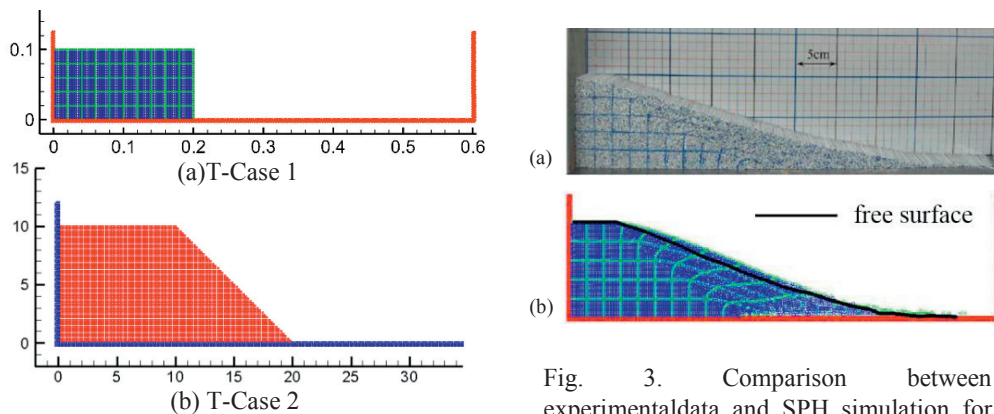


Fig.2. Initial state of test cases:
(a) T-Case 1, (b) T-Case 2

Fig. 3. Comparison between experimental data and SPH simulation for T-Case 1: (a) experimental result^[1]; (b) our simulated result

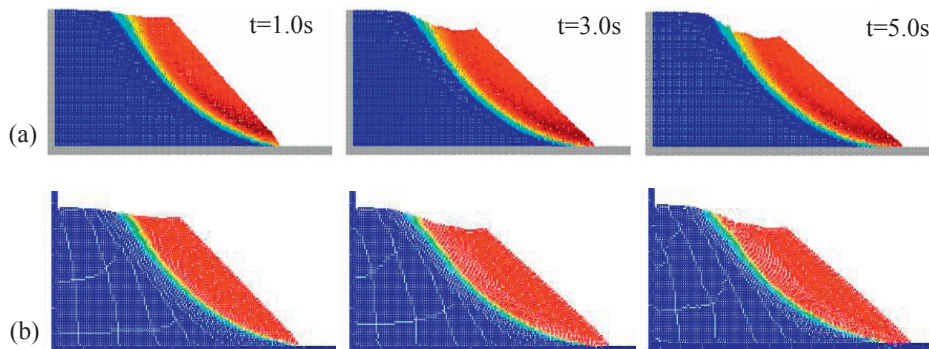


Fig.4 Contour plot of total displacement in the process of slope failure for T-Case 2:
a) SPH simulated result by Bui et al.^[8]; (b) our simulated result

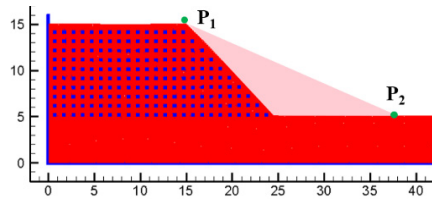


Fig.5. Model of earth-rock aggregate slope

Table 2. Specific parameter settings

| No. | Parameters | Soil | Rock |
|-----|------------------|-----------------------|-----------------------|
| 1 | Density | 1800kg/m ³ | 1800kg/m ³ |
| 2 | Elastic modulus | 15MPa | 50GPa |
| 3 | Poisson ratio | 0.25 | 0.25 |
| 4 | Cohesive force | 11kPa | / |
| 5 | Friction angle | 19.8° | / |
| 6 | Dilatancy angle | 0° | / |
| 7 | Kinetic friction | / | 0.65 |

Table 3. Summary of numerical cases

| Schemes \ Slope (°) | 25.0 | 27.5 | 30.0 | 32.5 | 35.0 | 37.5 | 40.0 | 42.5 | 45.0 |
|------------------------------------|----------|----------|----------|----------|----------|----------|----------|----------|----------|
| Cases without rocks | N-Case 1 | N-Case 2 | N-Case 3 | N-Case4 | N-Case 5 | N-Case 6 | N-Case 7 | N-Case 8 | N-Case 9 |
| Cases with small rocks (0.2m×0.2m) | S-Case 1 | S-Case 2 | S-Case 3 | S-Case 4 | S-Case 5 | S-Case 6 | S-Case 7 | S-Case8 | S-Case 9 |
| Cases with huge rocks (0.4m×0.4m) | H-Case 1 | H-Case 2 | H-Case 3 | H-Case4 | H-Case 5 | H-Case 6 | H-Case 7 | H-Case 8 | H-Case 9 |

4. Results and discussions

We employ this method to study the behavior of earth-rock aggregate. Fig.5 shows the numerical model. The blue blocks stand for the rocks and the red zone represents the soil. For simplicity, well-distribution of rocks is set up here. P_1 located at the slope top and P_2 located at the slope toe are the probe points. They respectively imply the horizontal and vertical deformations of the slope. Specific parameters are set as table 3. Numerical experiments consist of 3 schemes, respectively corresponding to case without rocks, case with small rocks (0.2m×0.2m) and case with huge rocks (0.4m×0.4m). The value range of slope angle is from 25-45 degree. It's worth mentioning that all of the cases with rocks feature the same blending ratio 15.6%. Fig.6 shows the process of slope failure with slope angle of 45° at representative times, via contour plots of velocity magnitude for (a) the case without rocks and (b) the case with huge rocks. It shows that the proposed soil-rock interaction algorithm works well in the SPH framework.

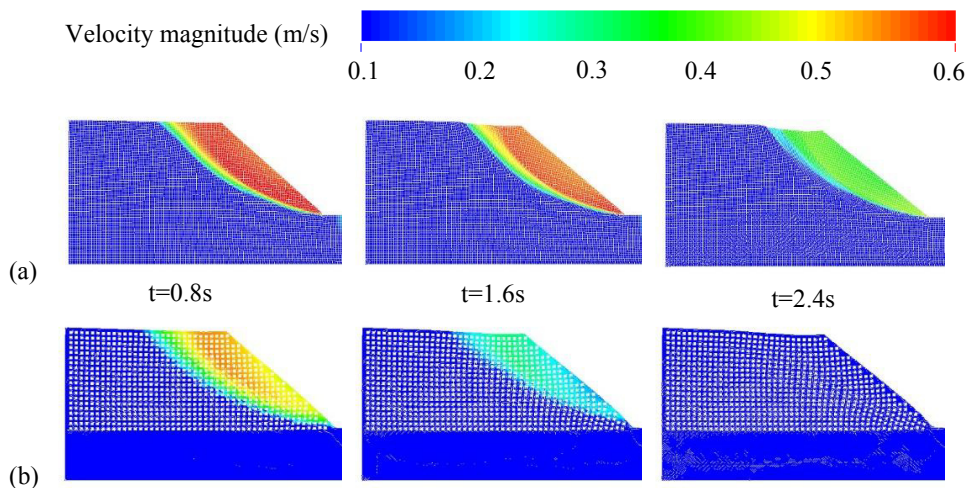


Fig.6 Contour plot of velocity magnitude in the process of slope failure with slope angle of 45°: (a) the case without rocks; (b) the case with small rocks

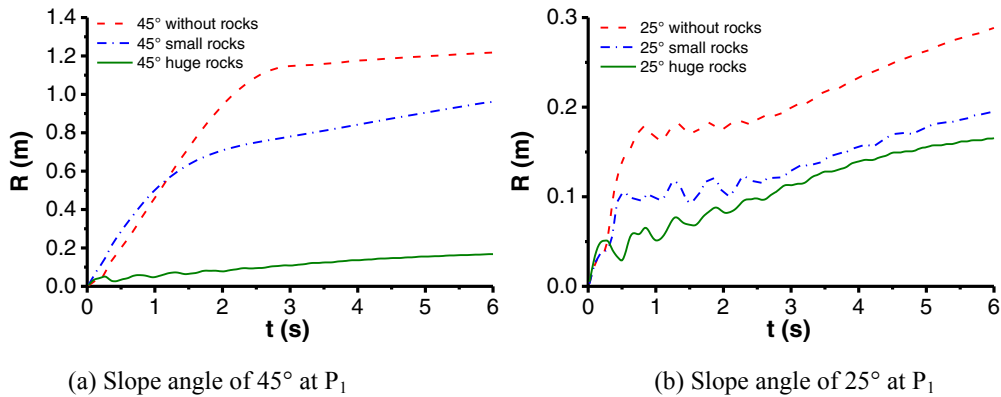


Fig.7. Comparison of the displacement-time curve at the slope top P_1 with (a) 25° and (b) 45° slope angle respectively for three schemes

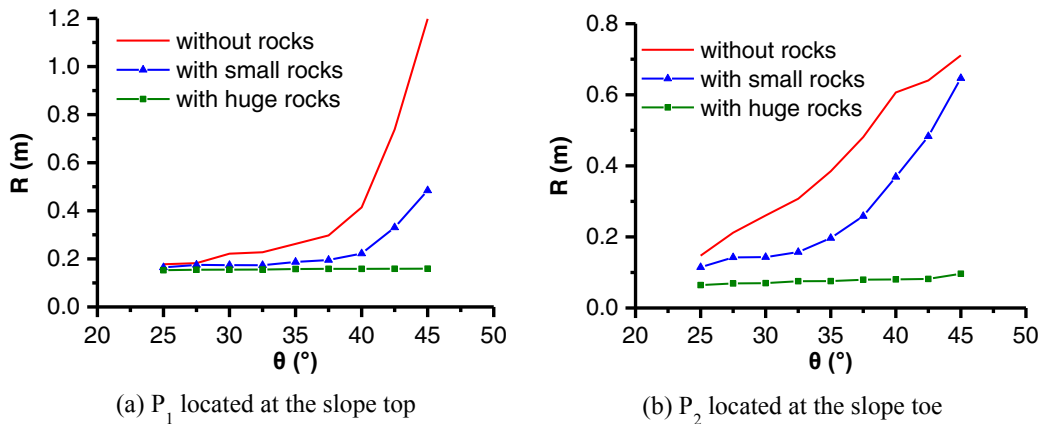


Fig.8. The relationship between displacement R and slope angle θ at (a) P_1 and (b) P_2 for three schemes at $t = 5.0$ s.

Fig.7 is the displacement-time graph at slope top P_1 for three schemes. According to Fig.7(a), when the slope angle is equal to 45°, the displacement value grows greatly over time, especially in the absence of rocks corresponding to the scheme 1. However, for scheme 3 (with huge rocks), the displacement is almost invariable all the time. And for scheme 2 (with relative small rocks), the time evolution of displacement falls in between. They show that the rocks would block the soil plastic flow. And with the same blending ratio, the blockade effect would be enlarged with the increase of rock size. When the slope angle is equal to 25°, as shown in Fig.7(b), the same phenomena is observed although the displacement varies weakly in either case. It partly suggests that the slope angle is another important factor that affects the whole deformation of landslide.

Let's focus on the typical time instant of 5.0 s and study the relationship between the displacement and the slope angle. Fig.8 shows the relationship between displacement R and slope angle θ at (a) P_1 and (b) P_2 for three schemes. As we know, with the increase of slope angle, the impact of soil resilience reduces gradually and so does the earth-rock aggregate. But even so, rocks could still block the development of landslide only if the rocks are big enough, as shown as the lower line for scheme 3 (with huge rocks) both in Fig.8(a) and Fig.8(b). For the other two schemes, the displacement grows greatly with the increase of the slope angle. Notably, the growing rate of displacement at P_1 is

widening especially within the range of $40^{\circ}\sim 45^{\circ}$. But an opposite state is found at P_2 in the same range. We deduce that the soil plastic flow will occur at P_2 but won't at P_1 . When the slope angle is too large, the blockade of rocks at P_2 is already not powerful enough to hold the soil plastic flow.

Conclusions

Based on elastic-plastic soil model, we proposed a soil-rock interaction algorithm, which allows to study more details in the post-failure of the earth-rock aggregate landslide. Numerical results obtained in this paper are qualitatively correct throughout. They showed that the rocks would block the development of landslide. With the same blending ratio, the blockade effect would be enlarged with the increase of stone size. The development of displacement at the slope toe is slightly with that at the slope top. The reason is just that plastic flow will occur at the slope toe but not at the slope top. And beyond that, the blockade effect of rocks is not apparent when the slope angle increasing to a certain degree.

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